PolarMonitoring project WP3: Simulation & performance analysis

WP31: New implementations in SMRT for altimetry applications

Progress meeting 2 - ESTEC - 4/10/2019



"To build a snow radiative transfer model to be integrated in the CLS altimeter simulation tool Altidop"

 Adapt SMRT to provide not only the total backscatter from the surface but also the time travel of the individual echo required by the simulator to re-construct the radar waveform

2) Definition of two study cases with different snow parameters



Coupling of a snow model with an altimetric model

1) Simulate **the total backscattering** with a Snow Radiative Transfer model: $\sigma_{TOT}(f, \theta_i) = \sigma_{surf}(f, \theta_i) + \sigma_{vol}(f, \theta_i)$

2) Estimate the propagation speed in each layer to introduce the time-dependency and rebuild the total echo



Theory to estimate the energy backscattered

With P_{layers} : surface echo, P_{grains}: volume echo

$$P_{\rm FV}(t) = \frac{1}{\left(4\pi\right)^3 H^4} \left[P_{\rm layers} + P_{\rm grains} \right]$$

 $P_{\text{layers}}(t) = \sum_{n=1}^{\infty} \lambda_n^2 \cdot A(z_n) \cdot \delta(t - t_v(z = z_n)) \qquad \otimes \int_S \delta(t - t'(z = 0, \theta)) \sigma_1^0(z = z_n, \theta) G^2(\theta) dS$ $P_{\text{grains}}(t) = \int_z \lambda^2(z) \cdot A(z) \cdot \delta(t - t_v(z)) \gamma_{\text{grains}}(z) dz \qquad \otimes \int_S \delta(t - t'(z = 0, \theta)) G^2(\theta) dS$

Vertical scattering distributions for the layers and the grains: SMRT outputs P_{FS}: SMRT P_{FS}: AltiDop

SMRT outputs:

- 1-t_v(z) with a vertical permittivity profile => new rt solver
- 2-Scattering of snow grains $\gamma_{grains}(z)$ with any formulation available in SMRT
- 3-A(z) is the power extinction at the depth z
- 4- Surface backscattering coefficient σ_l^0 profile given by SMRT
- => with an adapted surface model according to roughness heights.

Adaptation of SMRT for radar altimetry

New implementation: to output Backscatter contributions of every layer + Propagation speed of electromagnetic waves in each layer



For each many box formulations are available and new ones can be easily developed

IGE

Previous version : only "DORT" solver Now: « nadir altimetry » solver => to have the timetravel of the wave in each layer (1st order time

Adaptation of SMRT for radar altimetry





do l'environnement.

Adaptation of SMRT for radar altimetry

Different surface scattering models working with limited ranges of roughness: => Geometrical optics: Surface height roughness >> λ (>> 10 cm in S, >> 3 cm in Ku, >> 1cm in Ka) => IEM: Surface height rms < 10 cm in Ku band, and 5 cm in Ka



band

New implementations in SMRT

- Coded in Python

- Very easy to use, open-source and fully documented



Intrinsic validation of the total backscatter

Comparison between the two solvers: simulations with 2 simple 1-layer snowpack, at the nadir

Small surface height rms

Large surface height rms (more scatterings)



=> Good agreements between the two RT solvers at the nadir at low frequency => Show the effects of the 2nd order (multiple scattering) at high frequencies

Validation (Ku band): Comparison with Lacroix et al., 2008

Lacroix et al., 2008: altimeter model adapted for dual-frequecy, accounting for snowpack properties (simple sowpack) Comparison of waveform simulations with different snow grain sizes

A 1-layer snowpack with small snow grains

A 1-layer snowpack with large snow grains

120



- => Good agreement between volume echos.
- => Small differences between surface echos => work in progress in SMRT

Validation (S band): Comparison with Lacroix et al., 2008

Comparison of waveform simulations with different snow grain sizes

A 1-layer snowpack with small snow grains



A 1-layer snowpack with large snow grains

- => Good agreements between surface echos.
- => Small differences between volume echos for snowpack with large grains (numerical)

Validation: simple modeling of the Antarctic ice sheet

=> 2 realistic synthetic snowpack in Antarctica

Table 3

Snow parameters inversion at 2 locations (Vos=Vostok, Mau=Dronning Maud Land) based on the fit on individual waveforms represented in the Figs. 12 and 13

Location (lon, lat)	ρ_0 (g cm ⁻³)	μ (kg/m ²)	$\Phi_{\rm g}~({\rm mm})$	$\sigma_{\rm h}~({\rm mm})$	l (cm)
Vos (104.13, -76.57)	0.24±0.03	80±70	0.9 ± 0.07	3.9±0.3	14±3.6
Mau (9.83, -74.96)	$0.34 {\pm} 0.05$	422±339	$0.8\!\pm\!0.17$	$4.8\!\pm\!0.5$	33.5 ± 4.9
The error range is found by considering a signal to noise ratio on the waveforms of 10%.					

Say 154 Say 154 Anothor MR Lacroix et al., 2008



Validation: simple modeling of the Antarctic ice sheet

Ku Band (~13 GHz)

Vostok Lake ($\rho_0 = 240 \text{ kg m}^{-3}$, grain radius = 900 nm)

Dronning Maud Land ($\rho_0 = 340 \text{ kg m}^{-3}$, grain radius = 800 nm)

100

120



Validation: simple modeling of the Antarctic ice sheet

S Band (~3 GHz)

```
Vostok Lake
(rms = 3.9 mm, l = 4.8 cm, grain radius = 900 nm)
```



Dronning Maud Land (rms = 4.8 mm, l = 33 cm, grain radius = 800 nm)



Preliminary sensitivity analysis: Ku band



Preliminary sensitivity analysis: S band

Table 2 Range of variation of the snow parameters over the Antarctic ice-sheet 3.0 Te-23 3.0^{1e-23} Parameter Notation Min Max Step References --- SMRT - total volume - SMRT variations with temperature SMRT variations with surface RMS (units) value value --- SMRT - volume surface - SMRT variations with temperature 1 0.5 0.02 Goodwin et al. Surface 0.2 ρ_0 2.5 --- SMRT - surface total - SMRT variations with temperature 2.5 $(g \text{ cm}^{-3})$ density (1988)Accumulation 0 2000 Vaughan et al. μ - constant 2.0 2.0 $(kg m^{-2} y^{-1})$ (1999)rate $\Phi_{\rm g}$ (mm) 0.1 0.1Surdyk and Fily Snow grain 1 Java 1.5 Power Power (1993)size 20 Surface rms $\sigma_{\rm h}$ (mm) Lacroix et al. height (submitted for publication) 1.0 1.0 50 Correlation l(cm)4 2 Lacroix et al. length (submitted for publication) 0.5 0.5 $T(\mathbf{K})$ 190 260 10 King and Turner Snow temperature (1987)0.0 0.0 20 40 60 80 100 120 20 40 60 80 100 120 n Lacroix et al., 2008 gate gate 3.0 Te-23 3.0 Te-23 SMRT variations with surface density SMRT variations with snow grain size 2.5 2.5 Snowpack of reference: 2.0 2.0 $\rho_0 = 350 \text{kg m}^{-3}$, Lower 1.5 Power 1.2 grain radius =500 nm 1.01.0 Temperature = 225 K Thickness = 50 cm0.5 0.5 ~~~~~~~~~~~~~~~~~~~~~~ 00000000000000000 rms = 5 mm, l = 5 cm (IEM) 0000000000 0.0<u>↓</u>___0 0.000 0.0 120 20 60 80 100 20 60 80 100 120 n 40 40 gate gate

ons proscience

20

Conclusion

What has been done?

- New implementations in an existing snow radiative transfer model for altimetric applications,
- Intrinsic validations
- Eexternal validations with an existing altimeter simulator and realistic snowpack
- Preliminary sensitivity analysis with a simple snowpack

Work in progress

- Sensitivity analysis with study cases.
- Performance analysis:

Check the following hypothese => **Backscatter is computed assuming only first order scattering** OK for Ku-band, but has to be verify for Ka-band (interactions of order 2+ are less negligible...)

Perspective

Major difficulties: validation with in situ data. Work in progress at IGE...

=> new concept using passive/active microwave simulations to build a realistic study case. => using in-situ measurements of snow properties and surface roughness acquired during the Asuma traverse (2015/2016) in Antarctica => Sentinel-3 observations

Thank you for your attention



Validation (Ku band): Comparison with Lacroix et al., 2008

Lacroix et al., 2008: altimeter model adapted for dual-frequecy, accounting for snowpack properties (simple sowpack) Comparison of simulated scattering coefficients: SMRT/Lacroix et al. 2008

- σ_s according to surface height rms (IEM model)





=> Good agreements, differences due to numerical resolutions (especially with large grains)

How does SMRT work?

Existing version: SMRT is plane-parallel multi-layer radiative transfer model (time independent)



For each box many formulations are available and new ones can be easily developed

 T_B or σ_{TOT} at f and theta GE